

# Barchan dunes on bidispersed granular beds

C. A. Alvarez<sup>1,2</sup>, F. D. Cúñez<sup>1</sup>, E. M. Franklin<sup>1</sup>

<sup>1</sup>School of Mechanical Engineering, UNICAMP - University of Campinas,  
Rua Mendeleev, 200, Campinas, SP, Brazil

<sup>2</sup>Department of Atmospheric and Oceanic Sciences, University of California, Los Angeles,  
Los Angeles, CA 90095-1565, USA

## Key Points:

- Bidispersed grains segregate over barchan dunes with one species covering most of its surface
- A transient line, transverse to the flow direction, forms and migrates over the dune surface while segregation occurs
- We propose that the transverse line results from a competition between fluid entrainment and easiness of rolling

---

Corresponding author: Erick M. Franklin, [erick.franklin@unicamp.br](mailto:erick.franklin@unicamp.br)

**Abstract**

Barchans are dunes of crescentic shape found on Earth, Mars and other celestial bodies, growing usually on polydispersed granular beds. In this Letter, we investigate experimentally the growth of subaqueous barchans consisting of bidispersed grains. By tracking barchans in a water channel, we found that the grain distribution within the dune changes with the employed pair, and that a transient stripe appears on the dune surface, transverse to the flow direction, which we propose to be the result of a competition between the strength of fluid entrainment and easiness of rolling for each grain type. Our results provide new insights into barchan structures found in other planetary environments.

**Plain Language Summary**

This paper is devoted to crescent-shaped dunes, known as barchans, consisting of bidispersed grains. Barchans are found on Earth, Mars and other celestial bodies, with roughly the same morphology but different scales. Taking advantage of the much smaller and faster scales of subaqueous barchans (decimeter and minute, respectively), we performed experiments in a water channel where we obtained the evolution of barchan morphology and the motion of individual grains over the barchan surface. We found surface structures not still reported for aeolian barchans, but similar to those found on Mars at much larger and slower scales (kilometer and millennium, respectively), and explain their formation based on the motion individual grains. Our results open new possibilities to explain surface structures observed on Mars and other planetary environments.

**1 Introduction**

Barchan dunes, or simply barchans, are crescent-shaped dunes with horns pointing downstream that are frequently found on Earth, Mars and other celestial bodies (Bagnold, 1941; Herrmann & Sauermann, 2000; Hersen, 2004; Elbelrhiti et al., 2005; Claudin & Andreotti, 2006; Parteli & Herrmann, 2007; Courrech du Pont, 2015). They are formed under one-directional flows and limited amount of available grains and, under these conditions, present a robust shape that arise in diverse environments with a large range of scales, going from the kilometer and millennium in the case of Martian dunes down to the centimeter and minute in the aquatic case (Hersen et al., 2002; Claudin & Andreotti, 2006).

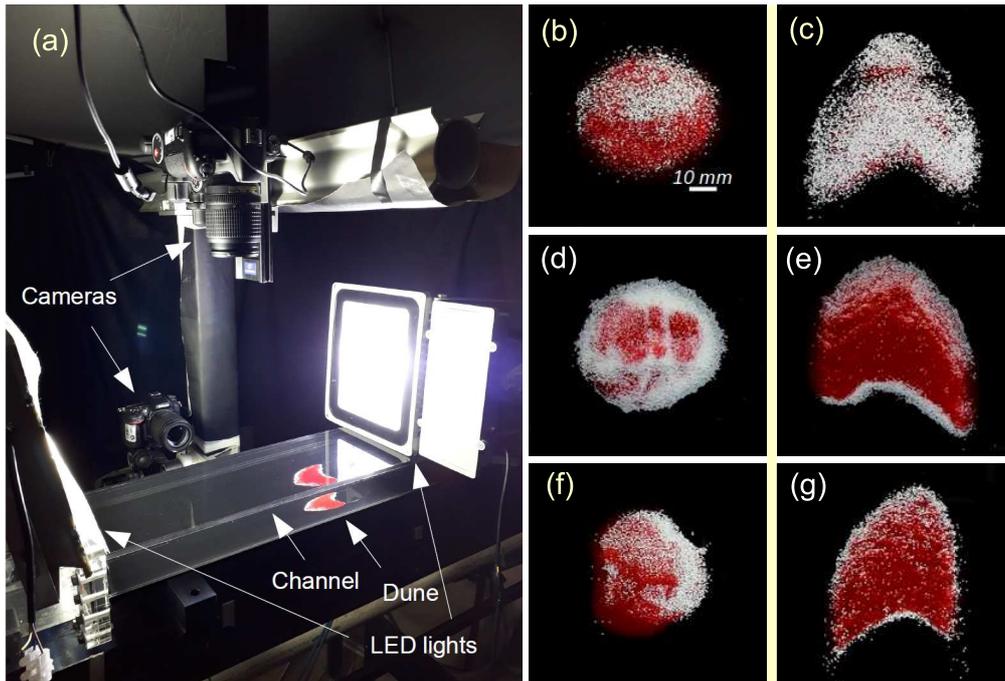
Because of their ubiquitous nature, barchans have been studied by using field measurements both on Earth (Finkel, 1959; Long & Sharp, 1964; Norris, 1966; Hesp & Hastings, 1998; Sauermann et al., 2000; Yang et al., 2019) and Mars (Breed et al., 1979; Schatz et al., 2006; Bishop, 2007; Bourke, 2010; Chojnacki et al., 2015; Runyon et al., 2017), which usually involve large time and length scales and uncontrolled conditions, producing valuable data that increased our understanding of many aspects of barchan morphodynamics, though with limited applicability for bedform evolution over long times and very few information at the grain scale. Given the smaller and faster scales of subaqueous barchans, experimental investigations were carried out in water channels and tanks under controlled conditions, from which it was possible to obtain the initial and long-time evolutions of the barchan morphology (Hersen et al., 2002; Alvarez & Franklin, 2017) and the motion of individual grains on the barchan surface (Alvarez & Franklin, 2018, 2019; Wenzel & Franklin, 2019). However, although natural beds are composed of polydispersed grains (Bagnold, 1941; Finkel, 1959; Long & Sharp, 1964), previous experiments investigated barchans of monodispersed particles, the only bidisperse-related references being the experiments of Caps and Vandewalle (2002) and Rousseaux et al. (2004) for size segregation within two-dimensional ripples and those of Groh et al. (2011) for density segregation within two-dimensional dunes.

Variations in the morphology and structure of barchans have been observed in nature, and in some cases have been described as resulting from hilly terrains (Finkel, 1959; Bourke, 2010; Parteli et al., 2014), wind changes (Finkel, 1959; Bourke, 2010; Parteli et al., 2014), barchan collisions (Long & Sharp, 1964; Hersen & Douady, 2005; Bourke, 2010; Vermeesch, 2011; Parteli et al., 2014; Assis & Franklin, 2020), or diverse processes over long time scales (Day & Catling, 2018). On Mars, barchan-shaped pits with stripes transverse to the flow direction were observed, which Day and Catling (2018) recently associated with dune casts formed by basalt flooding in ancient past and filled more recently with other sediments. Also on Mars, dunes with regions of different color, sometimes over the dune surface and other times on peripheral regions, were observed on its north pole and have been associated with the presence of coarse-grained ice together with sandy sediments (Breed et al., 1979; Schatz et al., 2006; Feldman et al., 2008; Ewing et al., 2010; C. J. Hansen et al., 2011; C. Hansen et al., 2013; Portyankina et al., 2013; Pommerol et al., 2013). For the latter dunes, the presence of two species with different sizes and densities could induce concentration of one of the species around the barchans, or induce segregation in horizontal layers. However, the precise organization of grains in those two-species barchans remains unknown.

In this Letter, we investigate experimentally the growth of subaqueous barchans consisting of bidispersed grains. We carried out exhaustive measurements where two-species barchans, in terms of grain sizes and/or densities, were tracked, and we varied the grain properties (size and density), their concentrations, and the water flow rate. We found that, while the barchan morphology remains roughly constant, denser, smaller, and smaller and less dense grains tend to accumulate on the barchan surface. For higher concentrations of the species accumulating on the surface, the other one forms a bottom layer and appears in top-view images as being around the barchan. We also found the appearance of a transient line on the dune surface, transverse to the flow direction, that forms during the growth of the single barchan, upstream its crest, and migrates toward the leading edge until disappearing. When grains of different densities but same size are used, that line consists of a large transverse stripe of different color. For the other pairs, that line consists in a transverse line that clearly separates a downstream region where segregation is complete from an upstream one where segregation is ongoing. The appearance of that transient line is intriguing, and we propose that it is the result of a competition between the strength of fluid entrainment and easiness of rolling for each grain type. For grains of different sizes, we observe, when the transient is finished, the presence of oblique stripes of much smaller wavelength than the dune size, which we associate with size segregation.

## 2 Materials and Methods

The experimental device consisted of a water reservoir, two centrifugal pumps, a flow straightener, a 5-m-long closed-conduit channel of transparent material and rectangular cross section (width = 160 mm and height  $2\delta = 50$  mm), a settling tank, and a return line. The last 2 m of the channel consisted of the 1-m-long test section followed by a 1-m-long section discharging in the settling tank. With the channel previously filled with water, mixtures of grains were poured inside, forming an initial pile of conical shape that was deformed afterward into a barchan dune by imposing a turbulent water flow. We used different grain sizes and densities that were mixed by pairs in different concentrations, and we varied the water flow rate. A camera placed above the channel acquired images of the bedforms, from which measurements at both the dune and grain scales were obtained by image processing (Alvarez & Franklin, 2017, 2018, 2019; Alvarez, 2020; Kelley & Ouellette, 2011; Crocker & Grier, 1996; Cúñez & Franklin, 2020), while a camera placed horizontally acquired side-view images used for measurements of the dune height. With this technique, we had access to the granular structure over the barchan surface, but not to its inner structure. Figure 1a shows a picture of the test section, and a lay-



**Figure 1.** (a) Photograph of the experimental setup showing the test section, cameras, LED lights and dune. (b) and (c) Top views of the initial pile and barchan dune, respectively, for case e of Table 1 (different densities and same diameter); (d) and (e) initial pile and barchan dune, respectively, for case h of Table 1 (same density and different diameters); (f) and (g) initial pile and barchan dune, respectively, for case q of Table 1 (different densities and diameters). In Figures (b) to (g), flow is from top to bottom.

115  
116

out of the experimental device and details about the used cameras are available in the supporting information.

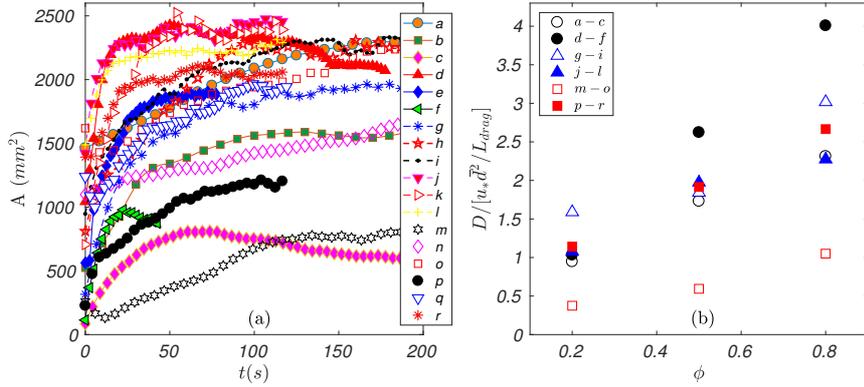
Case	$\phi_1$	$\phi_2$	$\phi_3$	red	Re	$d_{ratio}$	$\rho_{s,ratio}$	$m_0$
...	...	...	...	...	...	...	...	(g)
a	0.8	0.2	0	$S_2$	$1.47 \times 10^4$	1	1.64	15.5
b	0.5	0.5	0	$S_2$	$1.47 \times 10^4$	1	1.64	13.7
c	0.2	0.8	0	$S_2$	$1.47 \times 10^4$	1	1.64	11.8
d	0.8	0.2	0	$S_2$	$1.82 \times 10^4$	1	1.64	15.5
e	0.5	0.5	0	$S_2$	$1.82 \times 10^4$	1	1.64	13.7
f	0.2	0.8	0	$S_2$	$1.82 \times 10^4$	1	1.64	11.8
g	0	0.8	0.2	$S_3$	$1.47 \times 10^4$	2.5	1	10.5
h	0	0.5	0.5	$S_3$	$1.47 \times 10^4$	2.5	1	10.5
i	0	0.2	0.8	$S_3$	$1.47 \times 10^4$	2.5	1	10.5
j	0	0.8	0.2	$S_3$	$1.82 \times 10^4$	2.5	1	10.5
k	0	0.5	0.5	$S_3$	$1.82 \times 10^4$	2.5	1	10.5
l	0	0.2	0.8	$S_3$	$1.82 \times 10^4$	2.5	1	10.5
m	0.8	0	0.2	$S_3$	$1.47 \times 10^4$	2.5	1.64	15.5
n	0.5	0	0.5	$S_3$	$1.47 \times 10^4$	2.5	1.64	13.7
o	0.2	0	0.8	$S_3$	$1.47 \times 10^4$	2.5	1.64	11.8
p	0.8	0	0.2	$S_3$	$1.82 \times 10^4$	2.5	1.64	15.5
q	0.5	0	0.5	$S_3$	$1.82 \times 10^4$	2.5	1.64	13.7
r	0.2	0	0.8	$S_3$	$1.82 \times 10^4$	2.5	1.64	11.8

**Table 1.** Label of tested cases, initial concentration (volume basis) of each species within the initial pile, species with red (darker) color, channel Reynolds number  $Re$ , ratio between grain diameters  $d_{ratio}$ , ratio between grain densities  $\rho_{s,ratio}$ , and mass of the initial heap  $m_0$ .

117  
118  
119  
120  
121  
122  
123  
124  
125  
126  
127  
128  
129  
130  
131  
132  
133  
134  
135  
136  
137  
138  
139

The tests were performed with tap water at temperatures between 21 and 25 °C and three populations of particles were used: round zirconium beads (grain density  $\rho_s = 4100 \text{ kg/m}^3$ ) with grain diameters within  $0.40 \text{ mm} \leq d \leq 0.60 \text{ mm}$ , round glass beads ( $\rho_s = 2500 \text{ kg/m}^3$ ) with  $0.40 \text{ mm} \leq d \leq 0.60 \text{ mm}$ , and round glass beads with  $0.15 \text{ mm} \leq d \leq 0.25 \text{ mm}$ , which we call species 1, 2 and 3 ( $S_1$ ,  $S_2$  and  $S_3$ ), respectively, and for which  $\phi_1$ ,  $\phi_2$  and  $\phi_3$  are the used concentrations (see supporting information for microscopy images of the used grains). For each dune, grains of different species had either white or red colors in order to track their distribution over the barchan dune. The cross-sectional mean velocities of water  $U$  were 0.294 and 0.364 m/s, corresponding to Reynolds numbers based on the channel height,  $Re = \rho U 2\delta / \mu$ , of  $1.47 \times 10^4$  and  $1.82 \times 10^4$ , respectively, where  $\rho$  is the density and  $\mu$  the dynamic viscosity of the fluid. The shear velocities on the channel walls  $u_*$  were computed from velocity profiles measured with a two-dimensional particle image velocimetry device (2D-PIV), and were found to follow the Blasius correlation (Schlichting, 2000). By using the hydraulic diameter of the channel,  $u_*$  is found to be 0.0168 and 0.0202 m/s for the two flow rates employed, corresponding to variations of Reynolds numbers at the grain scale,  $Re_* = \rho u_* d / \mu$ , within 3 and 10 and to Shields numbers,  $\theta = (\rho u_*^2) / ((\rho_s - \rho)gd)$ , within 0.02 and 0.14, where we considered the midrange mean of the diameter of each species and  $g$  is the acceleration of gravity. All initial heaps had a volume of  $7.0 \text{ cm}^3$ , with initial masses varying between 10.5 and 15.5 g. The tested conditions are summarized in Table 1, where  $d_{ratio}$  and  $\rho_{s,ratio}$  are the ratios between grain diameters and densities, respectively (see supporting information for values of  $Re_*$  and  $\theta$  and pictures of initial and final bedforms for each test).

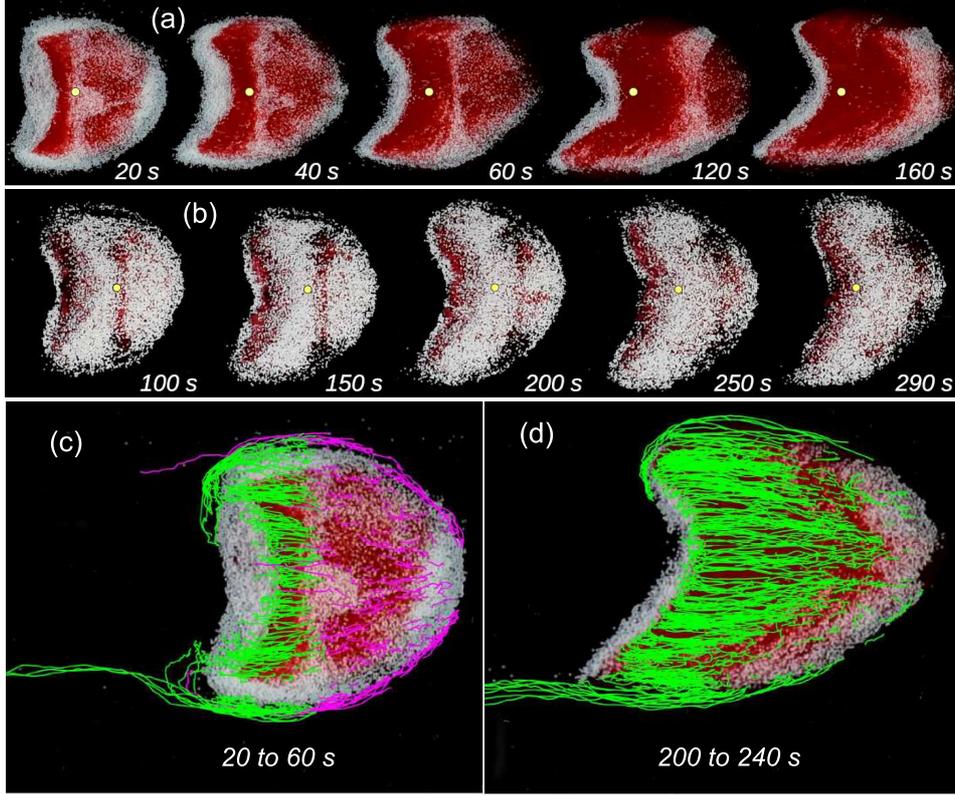
### 3 Results



**Figure 2.** (a) Projected area  $A$  of the dune occupied by the species accumulating over its surface as a function of time. Cases c and m had their time scale multiplied by 0.65 and 0.35, respectively, in order to fit the graphic. (b) Normalized diffusion-like coefficient  $D$  as a function of the concentration  $\phi$  of the species accumulating over the surface. Cases are listed in the key.

141 The general aspects of our results are summarized in Figures 1b to 1g. For different  
 142 mixtures, grains segregate forming two main layers, one over the other, where the  
 143 top layer consists of (i) denser grains when different densities are used (Figures 1b and  
 144 1c); (ii) smaller grains in the case of different diameters (Figures 1d and 1e); and (iii)  
 145 smaller and less dense grains for different diameters and densities (Figures 1f and 1g).  
 146 Those results are surprising, different gradings having been reported in the literature.  
 147 For different sizes, cases g to r of Table 1, measurements on two-dimensional subaqueous  
 148 ripples (Caps & Vandewalle, 2002; Rousseaux et al., 2004) showed that larger grains  
 149 tend to accumulate over their crest, while coarse sediments form a bottom layer that re-  
 150 mains immobile in the case of river dunes (Kleinhans, 2004). In the present case, although  
 151 forming a bottom layer, coarse grains are periodically entrained by the fluid flow. For  
 152 grains of different densities and same size, cases a to f of Table 1, a startling result had  
 153 already been found by Groh et al. (2011) for two-dimensional dunes, where denser par-  
 154 ticles were found to accumulate in the central body of the dune, close to its top but cov-  
 155 ered by a thin layer of less dense grains. An explanation for that grading covered by a  
 156 thin layer remains missing, and now we found something still different in the case of barchans.  
 157 In the present case, we propose, based on Makse (2000), that the observed grading is the  
 158 result of a competition between grains more easily entrained by the fluid flow and those  
 159 that roll more easily over the bed.

160 In order to evaluate the segregation in horizontal layers, we analyzed the increase  
 161 in the area occupied by grains accumulating over the dune surface. Figure 2a shows the  
 162 projected area  $A$  of the dune (top view) occupied by the species accumulating over its  
 163 surface as a function of time, for all the tested conditions, and it is a measure of the spread-  
 164 ing of those species over the surface. Figure 2a shows the existence of an initial phase, spread-  
 165 ing over the surface is occurring at a roughly constant rate, and a final  
 166 phase, reached when the quantity of grains of each species over the surface remains roughly  
 167 constant. Although grains are entrained by the fluid, we associate the spreading of grains  
 168 over the surface with a diffusion-like mechanism, in a similar way as in the turbulence  
 169 viscosity (Schlichting, 2000). For that, we defined a diffusion coefficient  $D$  ( $\text{mm}^2/\text{s}$ ), given  
 170 by the slope of initial phases of each curve in Figure 2a, and which we expect to scale,  
 171 as shown in Eq. 1, with the inertial length for the stabilization of sand flux (Hersen et



**Figure 3.** Evolution of the transverse stripe and trajectories of grains. (a) and (b) Snapshots of barchans at different instants showing the transverse stripe as it migrates to the leading edge. They correspond to cases h and b of Table 1, respectively, and the yellow circle indicates the crest position. (c) and (d) Trajectories of larger grains of case h, for two different intervals: when the stripe was still present and when it had already vanished, respectively. Trajectories were computed over 40 s and are superposed with the bedform image at the initial time of their computation. In Figure (c), green lines correspond to grains that started moving in positions downstream the transverse stripe, while magenta lines correspond to those that started moving in positions upstream the stripe.

172 al., 2002)  $L_{drag} = (\bar{\rho}/\rho)\bar{d}$ , the shear velocity  $u_*$ , and the mean diameter weighted by con-  
 173 centrations  $\bar{d} = \left(\frac{\phi_a}{d_a} + \frac{\phi_b}{d_b}\right)^{-1}$ , where  $\bar{\rho} = \phi_a\rho_a + \phi_b\rho_b$ , and  $a$  and  $b$  correspond to each  
 174 species. Figure 2b presents  $D(u_*\bar{d}^2/L_{drag})^{-1}$  as a function of particle concentration for  
 175 all tested cases, for which we obtain diffusion-like coefficients of the same order of mag-  
 176 nitude, around unity. Although there is some variation, a diffusion-like coefficient seems  
 177 proper for predicting the transient duration for the spreading of a species over the dune  
 178 surface.

$$D \sim \frac{u_*\bar{d}^2}{L_{drag}} \quad (1)$$

179 The general structures observed over barchans after segregation has taken place  
 180 can be better understood from the transverse line that forms on the dune surface dur-  
 181 ing the growth of the single barchan and migrates toward the leading edge until disap-

182 peering. The evolution of the transverse stripe is shown in Figure 3 together with tra-  
 183 jectories of larger grains migrating over the dune. Despite showing only some trajecto-  
 184 ries in Figures 3c and 3d, they represent well the behavior of the whole set. During the  
 185 transient, we can observe a region where segregation has already occurred, downstream  
 186 the transverse line, where larger grains (in cases of bidispersed size) and lighter grains  
 187 (in the case of bidespersed density) flight over longer distances arriving directly on the  
 188 crest and lee face, while the same does not occur on the region upstream the transverse  
 189 line. In the latter region, segregation occurs faster close to the transverse line, where larger  
 190 (cases g to r) or lighter (cases a to f) grains hop to the already segregated surface, and  
 191 from there roll/flight easier over the smaller grains (in cases of bidispersed size). With  
 192 that process, while larger or lighter grains travel to the lee face and form a carpet for  
 193 the smaller ones, as proposed by Groh et al. (2011) in the 2D case, the line is displaced  
 194 toward the leading edge until segregation is complete. When grains of different densi-  
 195 ties but same size are used, that line consists of a large transverse stripe of different color,  
 196 and is similar to structures observed on Mars, though the proposed mechanisms and ori-  
 197 gins found in the literature are different (Day & Catling, 2018). In our experiments, the  
 198 origin of the transverse stripe is always close to the crest of the initial pile, downstream  
 199 of which segregation occurs fast due to the formation of a recirculation region and lee  
 200 face, and from that position it propagates upstream with respect to the dune.

201 Finally, we observed also the presence of oblique stripes with wavelengths scaling  
 202 with the grain diameter, as can be observed in the upstream region of the dune shown  
 203 in Figure 1e. Based on images of cases g, h, j and k, we computed the arithmetic mean  
 204 of wavelengths of oblique stripes,  $\lambda$ , 20 s after the transverse stripe had vanished, and  
 205 found values of the order of  $10d$  ( $12 \leq \lambda/d \leq 16$ , see supporting information for more  
 206 images showing the oblique stripes and a table with  $\lambda$  for each case measured). It is dif-  
 207 ficult to explain their formation without measuring the inner grading of barchans; how-  
 208 ever, we speculate that their origin is a size segregation similar to those found in avalanches  
 209 (Makse et al., 1998), but with the driving force in the direction of water streamlines.

## 210 4 Conclusions

211 In conclusion, we observed a general segregation with one particular species cov-  
 212 ering most of the dune surface. This segregation could be at the origin, together with  
 213 solidification and frosting (as proposed by C. J. Hansen et al. (2011); C. Hansen et al.  
 214 (2013); Portyankina et al. (2013); Pommerol et al. (2013)), of part of the structures ob-  
 215 served on barchans on the north pole of Mars, for which top-view images show that dunes  
 216 consisting of coarse-grained ice together with other sandy sediments present one species  
 217 sometimes over the dune surface and other times on peripheral regions. We observed also  
 218 a transient line on the dune surface, transverse to the flow direction, that migrates to-  
 219 ward the leading edge, and that, when grains of different densities but same size are used,  
 220 consists of a large transverse stripe. Similar stripes are found on barchan-shaped pits  
 221 observed on Mars, resulting, however, from a different process (Day & Catling, 2018).  
 222 We propose that the transverse line, as well as the segregation in horizontal layers, are  
 223 the result of a competition between the strength of fluid entrainment and easiness of rolling  
 224 for the involved species. These findings open new possibilities to explain surface struc-  
 225 tures observed on Mars and shed light on the barchan dynamics over polydispersed beds.

## 226 Acknowledgments

227 C. A. Alvarez is grateful to SENESCYT (Grant No. 2013-AR2Q2850) and to  
 228 CNPq (Grant No. 140773/2016-9), F. D. Cúñez is grateful to FAPESP (Grant No.  
 229 2016/18189-0), and E. M. Franklin is grateful to FAPESP (Grant No. 2018/14981-7)  
 230 and to CNPq (Grant No. 400284/2016-2) for the financial support provided. Data

231 supporting this work are available in the supporting information and in Mendeley  
 232 Data (<https://data.mendeley.com/datasets/z42c97sw4c>).

## 233 References

- 234 Alvarez, C. A. (2020). *Genesis and formation of subaqueous barchan dunes : from*  
 235 *a morphological characterization to a description at the grain scale.* Un-  
 236 published doctoral dissertation, University of Campinas. Retrieved from  
 237 <http://repositorio.unicamp.br/jspui/handle/REPOSIP/343638>
- 238 Alvarez, C. A., & Franklin, E. M. (2017, Dec). Birth of a subaqueous barchan  
 239 dune. *Phys. Rev. E*, *96*, 062906. Retrieved from [https://link.aps.org/doi/](https://link.aps.org/doi/10.1103/PhysRevE.96.062906)  
 240 [10.1103/PhysRevE.96.062906](https://link.aps.org/doi/10.1103/PhysRevE.96.062906) doi: 10.1103/PhysRevE.96.062906
- 241 Alvarez, C. A., & Franklin, E. M. (2018, Oct). Role of transverse displacements in  
 242 the formation of subaqueous barchan dunes. *Phys. Rev. Lett.*, *121*, 164503. Re-  
 243 trieved from <https://link.aps.org/doi/10.1103/PhysRevLett.121.164503>  
 244 doi: 10.1103/PhysRevLett.121.164503
- 245 Alvarez, C. A., & Franklin, E. M. (2019, Oct). Horns of subaqueous barchan  
 246 dunes: A study at the grain scale. *Phys. Rev. E*, *100*, 042904. Retrieved  
 247 from <https://link.aps.org/doi/10.1103/PhysRevE.100.042904> doi:  
 248 [10.1103/PhysRevE.100.042904](https://link.aps.org/doi/10.1103/PhysRevE.100.042904)
- 249 Assis, W. R., & Franklin, E. M. (2020). A comprehensive picture for binary interac-  
 250 tions of subaqueous barchans. *Geophys. Res. Lett.*, *47*(18), e2020GL089464.
- 251 Bagnold, R. A. (1941). *The physics of blown sand and desert dunes.* London: Chap-  
 252 man and Hall.
- 253 Bishop, M. A. (2007). Point pattern analysis of north polar crescentic dunes, mars:  
 254 A geography of dune self-organization. *Icarus*, *191*(1), 151 - 157.
- 255 Bourke, M. C. (2010). Barchan dune asymmetry: Observations from Mars and  
 256 Earth. *Icarus*, *205*(1), 183 - 197.
- 257 Breed, C. S., Grolier, M. J., & McCauley, J. F. (1979). Morphology and distribu-  
 258 tion of common sand dunes on Mars: Comparison with the Earth. *J. Geophys.*  
 259 *Res.-Sol. Ea.*, *84*(B14), 8183-8204.
- 260 Caps, H., & Vandewalle, N. (2002). Patterns in hydraulic ripples with binary granu-  
 261 lar mixtures. *Physica A*, *313*(3), 357-364.
- 262 Chojnacki, M., Johnson, J. R., Moersch, J. E., Fenton, L. K., Michaels, T. I., & Bell,  
 263 J. F. (2015). Persistent aeolian activity at endeavour crater, meridiani planum,  
 264 mars; new observations from orbit and the surface. *Icarus*, *251*, 275 - 290.
- 265 Claudin, P., & Andreotti, B. (2006). A scaling law for aeolian dunes on Mars,  
 266 Venus, Earth, and for subaqueous ripples. *Earth Plan. Sci. Lett.*, *252*, 20-44.
- 267 Courrech du Pont, S. (2015). Dune morphodynamics. *C. R. Phys.*, *16*(1), 118 - 138.
- 268 Crocker, J. C., & Grier, D. G. (1996). Methods of digital video microscopy for  
 269 colloidal studies. *J. Colloid Interf. Sci.*, *179*(1), 298 - 310. Retrieved from  
 270 <http://www.sciencedirect.com/science/article/pii/S0021979796902179>  
 271 doi: <https://doi.org/10.1006/jcis.1996.0217>
- 272 Cúñez, F. D., & Franklin, E. M. (2020). Crystallization and jamming in narrow flu-  
 273 idized beds. *Phys. Fluids*, *32*(8), 083303.
- 274 Day, M. D., & Catling, D. C. (2018). Dune casts preserved by partial burial: The  
 275 first identification of ghost dune pits on mars. *J. Geophys. Res.-Planet*, *123*(6),  
 276 1431-1448.
- 277 Elbelrhiti, H., Claudin, P., & Andreotti, B. (2005). Field evidence for surface-wave-  
 278 induced instability of sand dunes. *Nature*, *437*(04058).
- 279 Ewing, R. C., Peyret, A.-P. B., Kocurek, G., & Bourke, M. (2010). Dune field pat-  
 280 tern formation and recent transporting winds in the olympia undae dune field,  
 281 north polar region of mars. *J. Geophys. Res.-Planet*, *115*(E8).
- 282 Feldman, W., Bourke, M., Elphic, R., Maurice, S., Bandfield, J., Prettyman, T., ...  
 283 Lawrence, D. (2008). Hydrogen content of sand dunes within Olympia Undae.

- 284 *Icarus*, 196(2), 422-432.
- 285 Finkel, H. J. (1959). The barchans of southern peru. *J. Geol.*, 67(6), 614-647.
- 286 Groh, C., Rehberg, I., & Kruehle, C. A. (2011, Nov). Observation of density seg-  
 287regation inside migrating dunes. *Phys. Rev. E*, 84, 050301(R). Retrieved  
 288from <https://link.aps.org/doi/10.1103/PhysRevE.84.050301> doi:  
 28910.1103/PhysRevE.84.050301
- 290 Hansen, C., Byrne, S., Portyankina, G., Bourke, M., Dundas, C., McEwen, A., ...  
 291Thomas, N. (2013). Observations of the northern seasonal polar cap on Mars:  
 292I. Spring sublimation activity and processes. *Icarus*, 225(2), 881-897.
- 293 Hansen, C. J., Bourke, M., Bridges, N. T., Byrne, S., Colon, C., Diniega, S., ...  
 294Thomas, N. (2011). Seasonal erosion and restoration of Mars' northern polar  
 295dunes. *Science*, 331(6017), 575-578.
- 296 Herrmann, H. J., & Sauermann, G. (2000). The shape of dunes. *Physica A (Amster-*  
 297*dam)*, 283, 24-30.
- 298 Hersen, P. (2004). On the crescentic shape of barchan dunes. *Eur. Phys. J. B*,  
 29937(4), 507-514.
- 300 Hersen, P., & Douady, S. (2005). Collision of barchan dunes as a mechanism of size  
 301regulation. *Geophys. Res. Lett.*, 32(21).
- 302 Hersen, P., Douady, S., & Andreotti, B. (2002, Dec). Relevant length  
 303scale of barchan dunes. *Phys. Rev. Lett.*, 89, 264301. Retrieved from  
 304<https://link.aps.org/doi/10.1103/PhysRevLett.89.264301> doi:  
 30510.1103/PhysRevLett.89.264301
- 306 Hesp, P., & Hastings, K. (1998). Width, height and slope relationships and aerody-  
 307namic maintenance of barchans. *Geomorphology*, 22, 193-204.
- 308 Kelley, D. H., & Ouellette, N. T. (2011). Using particle tracking to measure flow  
 309instabilities in an undergraduate laboratory experiment. *Am. J. Phys.*, 79(3),  
 310267-273.
- 311 Kleinhans, M. (2004). Sorting in grain flows at the lee side of dunes. *Earth-Science*  
 312*Reviews*, 65(1), 75-102.
- 313 Long, J., & Sharp, R. (1964). Barchan-dune movement in imperiel valley, california.  
 314*Bull. Geol. Soc. Am.*, 75, 149-156.
- 315 Makse, H. A. (2000). Grain segregation mechanism in aeolian sand ripples. *Eur.*  
 316*Phys. J. E*, 1, 127-135.
- 317 Makse, H. A., Havlin, S., King, P. R., & Stanley, H. E. (1998). Experimental studies  
 318of stratification in a granular heleshaw cell. *Philos. Mag. B*, 77(5), 1341-1351.
- 319 Norris, R. M. (1966). Barchan dunes of imperial valley, california. *J. Geol.*, 74(3),  
 320292-306.
- 321 Parteli, E. J. R., Durán, O., Bourke, M. C., Tsoar, H., Pöschel, T., & Herrmann,  
 322H. (2014). Origins of barchan dune asymmetry: Insights from numerical  
 323simulations. *Aeol. Res.*, 12, 121-133.
- 324 Parteli, E. J. R., & Herrmann, H. J. (2007, Oct). Dune formation on the present  
 325mars. *Phys. Rev. E*, 76, 041307. Retrieved from <https://link.aps.org/doi/10.1103/PhysRevE.76.041307> doi: 10.1103/PhysRevE.76.041307
- 326  
 327 Pommerol, A., Appr, T., Portyankina, G., Aye, K.-M., Thomas, N., & Hansen,  
 328C. (2013). Observations of the northern seasonal polar cap on Mars iii:  
 329CRISM/HiRISE observations of spring sublimation. *Icarus*, 225(2), 911-922.
- 330 Portyankina, G., Pommerol, A., Aye, K.-M., Hansen, C. J., & Thomas, N. (2013).  
 331Observations of the northern seasonal polar cap on Mars ii: HiRISE photo-  
 332metric analysis of evolution of northern polar dunes in spring. *Icarus*, 225(2),  
 333898-910.
- 334 Rousseaux, G., Caps, H., & Wesfreid, J. (2004). Granular size segregation in under-  
 335water sand ripples. *Eur. Phys. J. E*, 13, 213-219.
- 336 Runyon, K., Bridges, N., Ayoub, F., Newman, C., & Quade, J. (2017). An inte-  
 337grated model for dune morphology and sand fluxes on mars. *Earth Plan. Sci.*  
 338*Lett.*, 457, 204 - 212.

- 339 Sauermann, C., Rognon, P., Poliakov, A., & Herrmann, H. J. (2000). The shape of  
 340 the barchan dunes of Southern Morocco. *Geomorphology*, *36*, 47-62.
- 341 Schatz, V., Tsoar, H., Edgett, K. S., Parteli, E. J. R., & Herrmann, H. J. (2006).  
 342 Evidence for indurated sand dunes in the martian north polar region. *J.*  
 343 *Geophys. Res-Planet*, *111*(E4).
- 344 Schlichting, H. (2000). *Boundary-layer theory*. New York: Springer.
- 345 Vermeesch, P. (2011). Solitary wave behavior in sand dunes observed from space.  
 346 *Geophys. Res. Lett.*, *38*(22).
- 347 Wenzel, J. L., & Franklin, E. M. (2019). Velocity fields and particle trajectories for  
 348 bed load over subaqueous barchan dunes. *Granular Matter*, *21*, 321-334.
- 349 Yang, J., Dong, Z., Liu, Z., Shi, W., Chen, G., Shao, T., & Zeng, H. (2019). Migra-  
 350 tion of barchan dunes in the western quruq desert, northwestern china. *Earth*  
 351 *Surf. Process. Landforms*, *44*(10), 2016-2029.