Partitioning uncertainty in projections of Arctic sea ice

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Abstract. Improved knowledge of the contributing sources of uncertainty in projections of Arctic sea ice over the 21st century is essential for evaluating impacts of a changing Arctic ecosystem. Here, we consider the role of internal variability, model structure and emissions scenario in projections of Arctic sea-ice extent (SIE) by using six single model initial-condition large ensembles and a suite of models participating in Phase 5 of the Coupled Model Intercomparison Project. For projections of September Arctic SIE, internal variability accounts for as much as 60% of the total uncertainty in the next few decades, while emissions scenario dominates uncertainty toward the end of the century. Model structure accounts for approximately 70% of the total uncertainty by mid-century and declines to 20% at the end of the 21st century. For projections of wintertime Arctic SIE, internal variability contributes as much as 60% of the total uncertainty in the first few decades and impacts total uncertainty at longer lead times when compared to summer SIE. Model structure contributes the rest of the uncertainty with emissions scenario contributing little to the total uncertainty. At regional scales, the contribution of internal variability can vary widely and strongly depends on the month and region. For wintertime SIE in the GIN and Barents Seas, internal variability contributes approximately 70% to the total uncertainty over the coming decades and remains important much longer than in other regions. We further find that the relative contribution of internal variability to total uncertainty is state-dependent and increases as sea ice volume declines. These results demonstrate the need to improve the representation of internal variability of Arctic SIE in models, which is a significant source of uncertainty in future projections.

Keywords: *sea ice, climate change, uncertainty*

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1. Introduction

The rapid loss of Arctic sea ice over the last few decades has been one of the most iconic symbols of anthropogenic climate change. Since the beginning of the satellite record, September Arctic sea-ice extent (SIE) has decreased by approximately 50% (Stroeve and Notz, 2018) and experienced considerable thinning largely due to a lengthening of the melt season (Perovich and Polashenski, 2012; Stroeve et al., 2014). While state-of-the-art global climate models (GCMs) predict a decline of Arctic SIE throughout the 21st century, the exact amount of ice loss remains highly uncertain (Massonnet et al., 2012; SIMIP, 2020). Studies suggest that in the summertime the Arctic will most likely be "ice free" by the end of the 21st century (Jahn, 2018; Niederdrenk and Notz, 2018; Sigmond et al., 2018) and could possibly be ice free as early as 2050 (Jahn, 2018) or 2030 (Wang and Overland, 2009). To improve projections of Arctic sea ice, the relative importance of the sources of uncertainty need to be characterized and if possible reduced, particularly at regional scales (Eicken, 2013; Barnhart et al., 2016).

Internal variability, which refers to natural fluctuations in climate that occur even in the absence of external forcing, has long been known as an important source of uncertainty in projections of future climate (Hawkins and Sutton, 2009; Deser et al., 2012, 2020; Lehner et al., 2020; Maher et al., 2020). These fluctuations — intrinsic to the climate system — have been shown to exert a strong influence on short-term trends in numerous climate variables, such as surface temperature (Wallace et al., 2012; Smoliak et al., 2015; Deser et al., 2016; Lehner et al., 2017), precipitation (Hawkins and Sutton, 2011; Deser et al., 2012), snowpack (Siler et al., 2019), glacier mass balance (Marzeion et al., 2014; Bonan et al., 2019), ocean biogeochemical properties (Lovenduski et al., 2016; Schlunegger et al., 2020), and sea ice (Kay et al., 2011; Swart et al., 2015; Jahn et al., 2016; Screen and Deser, 2019; Rosenblum and Eisenman, 2017; England et al., 2019; Ding et al., 2019; Landrum and Holland, 2020). Recent estimates suggest that internal variability has contributed to approximately 50% of the observed trend in September Arctic SIE decline since 1979 (Stroeve et al., 2007; Kay et al., 2011; Zhang, 2015; Ding et al., 2017, 2019) and has strongly controlled regional patterns of sea ice loss (England et al., 2019).

The large role of internal variability in determining changes to Arctic SIE over the observational record means the predictability of future Arctic SIE at decadal timescales could remain heavily influenced by internal variability. The advent of decadal prediction systems (e.g., Meehl et al., 2009, 2014) raises the question whether realistic physics together with proper initialization of observations can lead GCMs to successfully constrain this internal variability and result in skillful estimates of SIE at decadal lead times (Koenigk et al., 2012; Yang et al., 2016). Initial-value predictability of Arctic SIE has been shown to be regionally and seasonally dependent (Blanchard-Wrigglesworth et al., 2011b; Bushuk et al., 2019), often only lasting a few years at most for total Arctic SIE

(Blanchard-Wrigglesworth et al., 2011a; Guemas et al., 2016). Using a suite of perfect model experiments (which quantify the upper limits of predictability), Yeager et al. (2015) showed that the rate of sea ice loss in the North Atlantic may slow down in the coming decades due to a reduction of ocean heat transport into the Arctic, which itself is highly predictable. Similarly, Koenigk et al. (2012) found a link between meridional overturning circulation and the potential predictability of decadal mean sea ice concentration in the North Atlantic — consistent with Yang et al. (2016). Indeed, this means that uncertainty due to internal variability is an important — and possibly reducible — source of uncertainty for short-term projections in some regions with properly initialized forecasts, but not for long-term projections. However, even if uncertainty due to internal variability cannot be reduced, understanding its magnitude will allow for better decision making in light of that uncertainty. This raises an important question: what is the relative role of internal variability in future projections of Arctic sea ice? Any accounting for the sources of uncertainty in projections of Arctic SIE must quantify the relative importance of each source at different spatial and temporal scales. For example, how important is internal variability for projections of Arctic sea ice 15 versus 30 years from now? Moreover, because models exhibit different magnitudes of internal variability in sea ice, particularly at regional scales (e.g., England et al., 2019; Topál et al., 2020), such quantification must sample the influence of model uncertainty in the estimate of internal variability itself.

To examine these questions we use an unprecedented suite of single model initial-condition large ensembles (SMILEs) from six fully-coupled GCMs. Due to their sample size, these SMILEs uniquely allow us to partition uncertainty in projections of Arctic SIE into the relative roles of internal variability, model structure, and emissions scenario at both Arctic-wide and regional spatial scales without relying on statistical representations of the forced response or internal variability (e.g., Lique et al., 2016). The SMILEs also allow us to quantify the influence of different estimates of internal variability, a feature of sea ice projection uncertainty that has received little attention. In what follows, we first investigate the role of internal variability in projections of total Arctic SIE. We then explore how the relative partitioning of each source changes as a function of season and Arctic region and how this partitioning is influenced by the mean-state of Arctic sea ice.

2. Data

2.1. MMLEA output

We use six SMILEs from the Multi-Model Large Ensemble Archive (MMLEA; Deser et al., 2020) to investigate the role of internal variability on projections of Arctic sea ice. These include the: 40 member Community Earth System Model Large Ensemble Community Project (CESM1-LE; Kay et al., 2015), 50 member Canadian

Earth System Model Large Ensemble (CanESM2-LE; Kirchmeier-Young et al., 2017), 30 member Commonwealth Scientific and Industrial Research Organisation Large Ensemble (CSIRO-Mk3.6.0-LE; Jeffrey et al., 2013), 20 member Geophysical Fluid Dynamics Laboratory Large Ensemble (GFDL-CM3-LE; Sun et al., 2018), 30 member Geophysical Fluid Dynamics Laboratory Earth System Model Large Ensemble (GFDL-ESM2M-LE; Rodgers et al., 2015), and 100 member Max Planck Institute Grand Ensemble (MPI-GE; Maher et al., 2019). Each SMILE uses historical and RCP8.5 forcing. We also use the RCP2.6 and RCP4.5 100 member ensembles from the MPI-GE. From each SMILE we use sea ice concentration (SIC) to compute monthly Arctic SIE (defined as the area where SIC>15%) for 6 Arctic regions and the pan-Arctic (see Figure S1). We also use sea ice thickness to compute monthly Arctic sea-ice volume (SIV) for these same spatial domains. Note that the output from GFDL-CM3 and GFDL-ESM2M is the average thickness over the ice-covered area of the grid cell. To compute SIV, the monthly averaged ice-covered thickness from both models was multiplied by the monthly average SIC of each cell to get the grid-cell average SIT. Prior to these calculations, all model output is regridded to a common $1^\circ \times 1^\circ$ analysis grid using nearest-neighbor interpolation.

2.2. CMIP5 output

We use monthly output from the historical, RCP2.6, RCP4.5, and RCP8.5 simulations of 30 different GCMs participating in CMIP5 (Taylor et al., 2012). Since the historical simulations end in 2005, we merge the 1850-2005 fields from the historical simulations with the 2006-2100 fields under each RCP forcing scenario. For each experiment, we use SIC to compute monthly Arctic SIE (defined as the area where SIC>15%). The set of GCMs evaluated reflects those that provide the necessary output (see Table S1). All model output is regridded to a common $1^\circ \times 1^\circ$ analysis grid using nearest-neighbor interpolation.

3. Uncertainty in projections of Arctic sea ice

We begin by partitioning three sources of uncertainty following Hawkins and Sutton (2009) and Lehner et al. (2020), where the total uncertainty (T) is the sum of the uncertainty due to model structure (M), the uncertainty due to internal variability (I) and the uncertainty due to emissions scenario (S). Each source can be estimated for a given time t and location x such that:

$$T(t,x) = I(t,x) + M(t,x) + S(t,x) \tag{1}$$

where the fractional uncertainty from a given source is calculated as I/T, M/T, and S/T. I is calculated as the variance across ensemble members of each SMILE, yielding one time-varying estimate of I per SMILE. Averaging across the six I yields the multi-model mean internal variability uncertainty (see white line in Figure 1). To quantify the influence of model uncertainty in the estimate of I we also use the model with the largest and smallest I (see white shaded regions in Figure 1). Model uncertainty in the estimate of I has emerged as an important and potentially reducible source of uncertainty in regional temperature and precipitation changes (Lehner et al., 2020; Deser et al., 2020) and projections of global ocean biogeochemical properties (Schlunegger et al., 2020). M is calculated as the variance across the ensemble means of the six SMILEs. It is important to note that the SMILEs used in this study are found to be reasonably representative of the CMIP5 inter-model spread for the percent of remaining Arctic sea ice cover (see Figure S2) and total Arctic SIE (see black lines in Fig. 1), but a more systematic comparison is necessary before generalizing this conclusion. Finally, since only a few of the SMILEs were run with more than one emissions scenario, we turn to CMIP5 for S, which is calculated as the variance across the multi-model mean RCP scenarios (see Table S1 for details). Prior to these variance calculations, the monthly SIE was smoothed with a 5-year running mean to isolate the effect of uncertainty on short-term projections and then used to calculate the percent of remaining sea ice relative to 1995-2014 (see Figure S2).

3.1. Total Arctic sea-ice extent

We first consider projections of Arctic SIE in September (the seasonal minimum) and March (the seasonal maximum). Figure 1 shows the fractional contribution of each source of uncertainty to total uncertainty. In September, uncertainty due to internal variability is important initially, accounting for approximately 30% of total uncertainty. However, over time model uncertainty increases and eventually dominates for the first half of the 21st century, before scenario uncertainty starts to dominate after approximately mid-century (Fig. 1c). However, model uncertainty in internal variability itself can have an effect on climate projections (e.g., Lehner et al., 2020). Accounting for the minimum and maximum contribution of internal variability to total uncertainty suggests that internal variability could account for as much as 50-60% or as little as 10-20% of total uncertainty in projections of September SIE in the coming decades and could contribute approximately 10% throughout the 21st century. Note, these results are similar for most summer months and summertime averages (see Fig. S4 and S5).

A different story emerges for projections of Arctic SIE in March. While uncertainty due to internal variability is again important initially and accounts for more of the total uncertainty at longer lead times, model uncertainty increases and quickly dominates until the end of the century (Fig. 1d). Scenario uncertainty is relatively less important for projections of Arctic SIE in March and, more broadly, during the wintertime (see Fig. S4). Uncertainty in model internal variability remains large throughout the 21st century, suggesting internal variability could account for as much as 20% or as little as 5% of the total uncertainty beyond mid-century. The relative partitioning is similar for most winter months and wintertime averages (see Fig. S4 and S5).

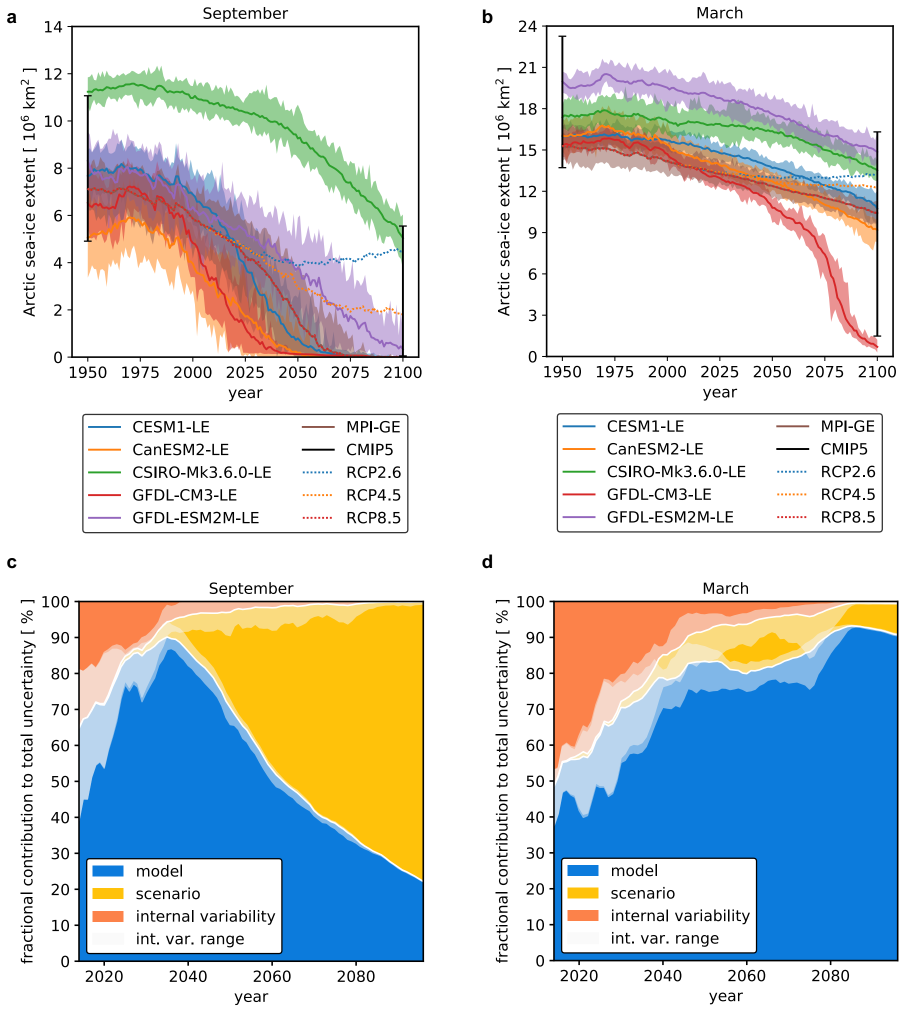

Figure 1. (a-b) Arctic sea-ice extent from 1950-2100 for six single model initial-condition large ensembles (SMILEs) in (a) September and (b) March. The bold line represents the ensemble-mean of each SMILE and the shading represents the range of each SMILE under historical and RCP8.5 forcing. The colored dotted lines represent the RCP scenarios from CMIP5 (shown only for the MPI-GE). The black vertical lines at 1950 and 2100 represent the spread from the 30 CMIP5 simulations. (c-d) Fractional contribution of model structure, emissions scenario, and internal variability to total uncertainty for the percent of remaining Arctic sea ice cover in (c) September and (d) March. The solid white lines denote the borders between each source of uncertainty, while the transparent white shading around those lines is the range of this estimate based on different estimates of internal variability in the MMLEA. Both fractional uncertainty panels are for five-year mean projections of percent of remaining Arctic sea-ice cover relative to 1995-2014.

These results suggest that uncertainty in short-term projections of Arctic sea ice, regardless of the season, is dominated by internal variability, while for long-term projections of Arctic sea ice, both scenario and model uncertainty become important. At long lead times, scenario uncertainty accounts for most of the uncertainty in projections of Arctic SIE in the summer months and model uncertainty accounts for most of the uncertainty in projections of Arctic SIE in the winter months. This likely reflects the fact that September Arctic SIE disappears in most GCMs by 2100 under RCP8.5.

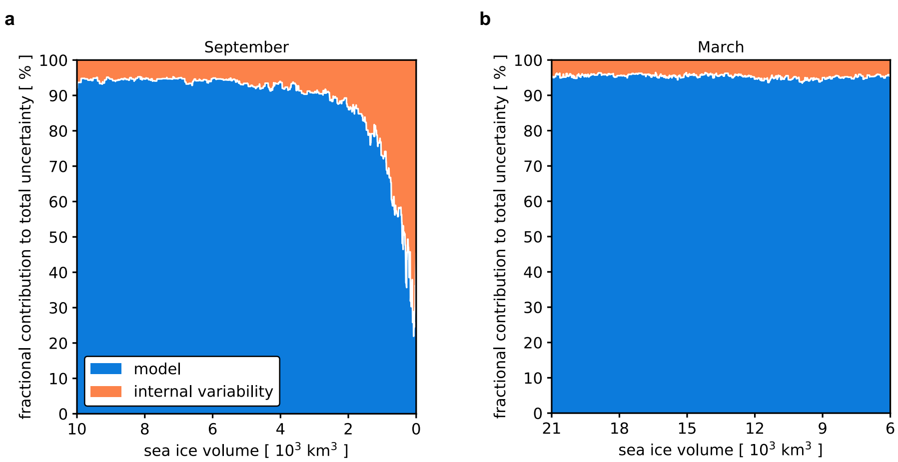

Figure 2. Fractional contribution of model structure and internal variability to total uncertainty for Arctic sea-ice extent (SIE) in (a) September and (b) March as a function of Arctic sea-ice volume (SIV). The solid white lines denotes the border between the two sources of uncertainty. Both fractional uncertainty panels are for projections of Arctic sea-ice extent with no temporal averaging or reference period. Note the x-axis is different for (a) and (b).

3.2. State dependence of internal variability

These results show a clear time-scale dependence for the relative importance of internal variability in uncertainty of projections of Arctic SIE. However, recent studies have shown that the internal variability and the predictability of Arctic sea ice can change over time and under anthropogenic forcing (Goosse et al., 2009; Mioduszewski et al., 2019; Holland et al., 2019). September Arctic SIE variability is expected to increase under warming (Goosse et al., 2009; Mioduszewski et al., 2019), suggesting that the role of internal variability in sea ice projections is mean-state dependent. To investigate the role of internal variability in projections of Arctic sea ice as a function of the mean-state, we partition the relative sources of uncertainty with respect to SIV by binning a given SIE to its associated SIV for each month. We then perform the same variance analysis described above as a function of SIV instead of as a function of time. Doing this for each SMILE member and the ensemble-mean of each SMILE allows us to examine the contributing sources of uncertainty as a function of SIV.

Figure 2 shows the fractional contribution of internal variability and model structure to total uncertainty for future Arctic SIE in September and March as a function of September and March Arctic SIV, respectively. Note, scenario uncertainty was excluded in these calculations (by using simulations from RCP 8.5 only) to isolate the effect of internal variability at different mean-states with respect to model uncertainty under the same mean-state. In September, as SIV declines — which is expected to occur throughout the 21st century — internal variability remains constant for most SIV values, accounting for approximately 10% of total uncertainty. However, at lower SIV regimes (< 3,000 km^3), the contribution of internal variability increases and accounts for approximately 80% of the total uncertainty at low thickness sea ice regimes (i.e., SIV < 1,000 km^3). This is consistent with previous work that has shown increased variability of summer Arctic SIE as it approaches zero (e.g., Mioduszewski et al., 2019). In March, the contribution of internal variability to total uncertainty remains relatively constant at all SIV regimes, likely reflecting the fact that sea ice is present in most winter climates in future projections (e.g., Goosse et al., 2009). It is important to note that this increase in the contribution of internal variability to uncertainty at lower SIV regimes holds for summer (June, July, and August) months (not shown).

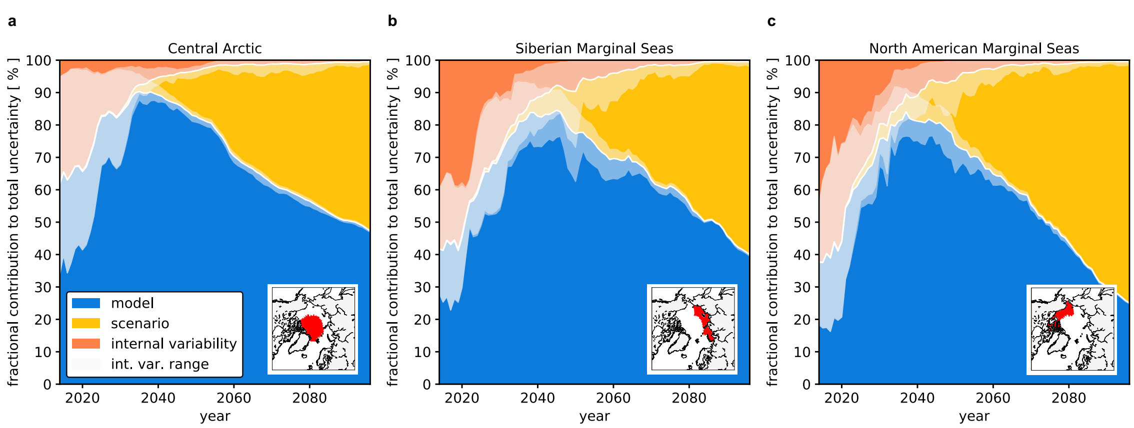

Figure 3. Fractional contribution of model structure, emissions scenario, and internal variability to total uncertainty for percent of remaining sea ice cover in July, August and September (JAS) for the Central Arctic, Siberian Marginal Seas (Kara Sea, Laptev Sea, East Siberian Sea), and North American Marginal Seas (Chukchi Sea, Beaufort Sea, Canadian Archipelago). The solid white lines indicate the borders between sources of uncertainty, while the transparent white shading around those lines is the range of this estimate based on different estimates of internal variability in the MMLEA. All panels are for five-year mean projections of percent of remaining Arctic sea-ice cover relative to 1995-2014.

3.3. Regional Arctic sea-ice extent

While the loss of total Arctic SIE is important for understanding the global climate response, climate change and sea ice loss are experienced predominately at regional scales (Barnhart et al., 2014; Lehner and Stocker, 2015). To investigate uncertainty in regional SIE projections, we compute SIE for 6 Arctic regions, which include the Central Arctic, Siberian Marginal Seas, North American Marginal Seas, Baffin/Hudson Bay and the Labrador Sea, the Bering Sea and Sea of Okhotsk, and Greenland-Iceland-Norwegian (GIN) and Bering Seas. These regions were chosen to represent geographically distinct parts of the Arctic ocean, where SIE retreat occurrs with different velocities. As with total Arctic SIE, the SMILEs used in this study are found to be reasonably representative of the CMIP5 inter-model spread for the percent of remaining Arctic sea ice cover in each region (see Figure S3).

Figure 3 shows the fractional contribution of each source of uncertainty to total uncertainty in projections of July, August, and September (JAS) SIE in the Central Arctic (Fig. 3a), Siberian Marginal Seas (Fig. 3b), and North American Marginal Seas (Fig. 3c). We only show summertime SIE as these regions are fully ice covered in the wintertime and exhibit little wintertime variability throughout much of the 21st century. As with total September Arctic SIE, there is a large role for internal variability initially, accounting for as much as 80% of total uncertainty in the Siberian and North American Marginal Seas (Fig. 3b and 3b) and 60% in the Central Arctic (Fig. 3a). However, over time model uncertainty increases and eventually dominates for the first half of the 21st century in Central Arctic (Fig. 3a) and marginal seas (Fig. 3b and Fig. 3c), accounting for 40-50% of the total uncertainty. Note, the contribution of model structure to total uncertainty at the end of the century is lowest for the North American Marginal Seas. By the end of the 21st century scenario uncertainty dominates and accounts for over half of the uncertainty, meaning that whether or not an ice free Arctic occurs in the summertime is a direct consequence of climate change policy. Notably, the range of simulated internal variability contributions remain quite large through the 21st century in each region.

Figure 4 shows the fractional contribution of each source of uncertainty to total uncertainty in projections of January, February, and March (JFM) Arctic SIE in Baffin Bay, Hudson Bay and the Labrador Sea (Fig. 4a), Bering Sea and Sea of Okhotsk (Fig. 4b), and GIN and Barents Seas (Fig. 4c). These regions were selected to examine wintertime SIE as there is highly variable SIE in winter and little-to-no SIE in summer. As with regions of variable summer sea ice cover, these regions show a distinct pattern of uncertainty partitioning. For Baffin Bay, Hudson Bay, and Labrador Sea, approximately 80% of total uncertainty in the next few decades is attributable to internal variability. Note that the contribution of uncertainty in the estimate of internal variability itself can cause this to change to only 20% (mainly driven by CSIRO-Mk3.6.0, which clearly

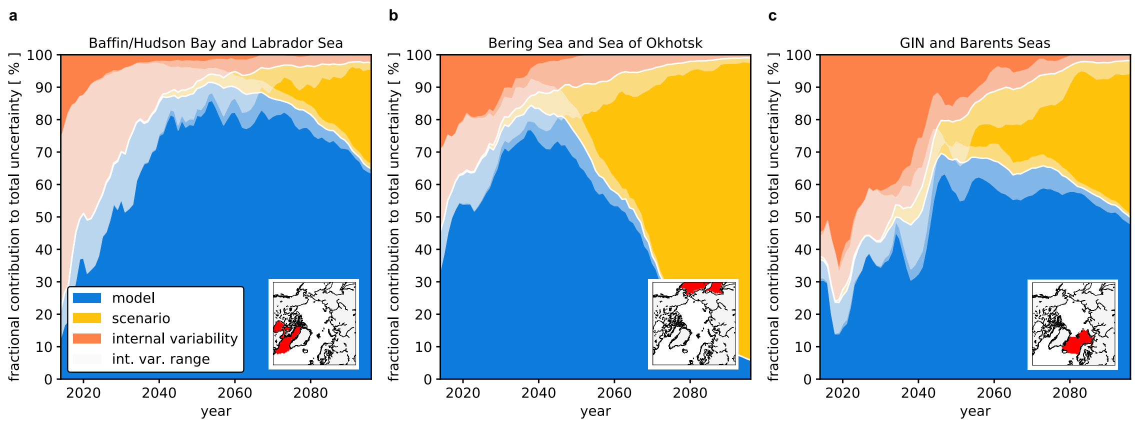

Figure 4. Fractional contribution of model structure, emissions scenario, and internal variability to total uncertainty for percent of remaining sea ice cover in January, February, and March (JFM) for (a) Baffin Bay, Hudson Bay, and the Labrador Sea, (b) Being Sea and Sea of Okhotsk, and the (c) GIN and Barents Seas. The solid white lines indicate the borders between sources of uncertainty, while the transparent white shading around those lines is the range of this estimate based on different estimates of internal variability in the MMLEA. All panels are for five-year mean projections of percent of remaining Arctic sea-ice cover relative to 1995-2014.

overestimates sea-ice extent in this region and Arctic-wide). The internal variability contribution diminishes to approximately 10% by the end of the century, and model structure dominates by 2030. A similar picture emerges for the Bering Sea and Sea of Okhotsk, but instead scenario uncertainty dominates in the latter half of the 21st century. Interestingly, the uncertainty partitioning for the GIN and Barents Seas has a distinct structure: internal variability dominates projection uncertainty for the next 30 years and remains persistent throughout much of the 21st century. The contribution of internal variability is notably larger than in other regions and is most likely related to the influence of Atlantic heat transport on sea ice (Årthun et al., 2012).

A key result here — in contrast to total Arctic SIE for March and September — is the larger role of internal variability in contributing to total uncertainty, which persists throughout much of the 21st century. This suggests decadal predictions of regional Arctic SIE will be highly influenced by internal variability, especially for wintertime conditions in the GIN and Barents Seas. Moreover, the range of internal variability across models presents a unique challenge as internal variability could account for as much as 80% or as little as 20% of the total uncertainty in regions like the Labrador Sea in the coming decades. Understanding the cause of the range in this internal variability uncertainty is an important next step, whether it is related to model biases or dependent on the sea ice mean-state.

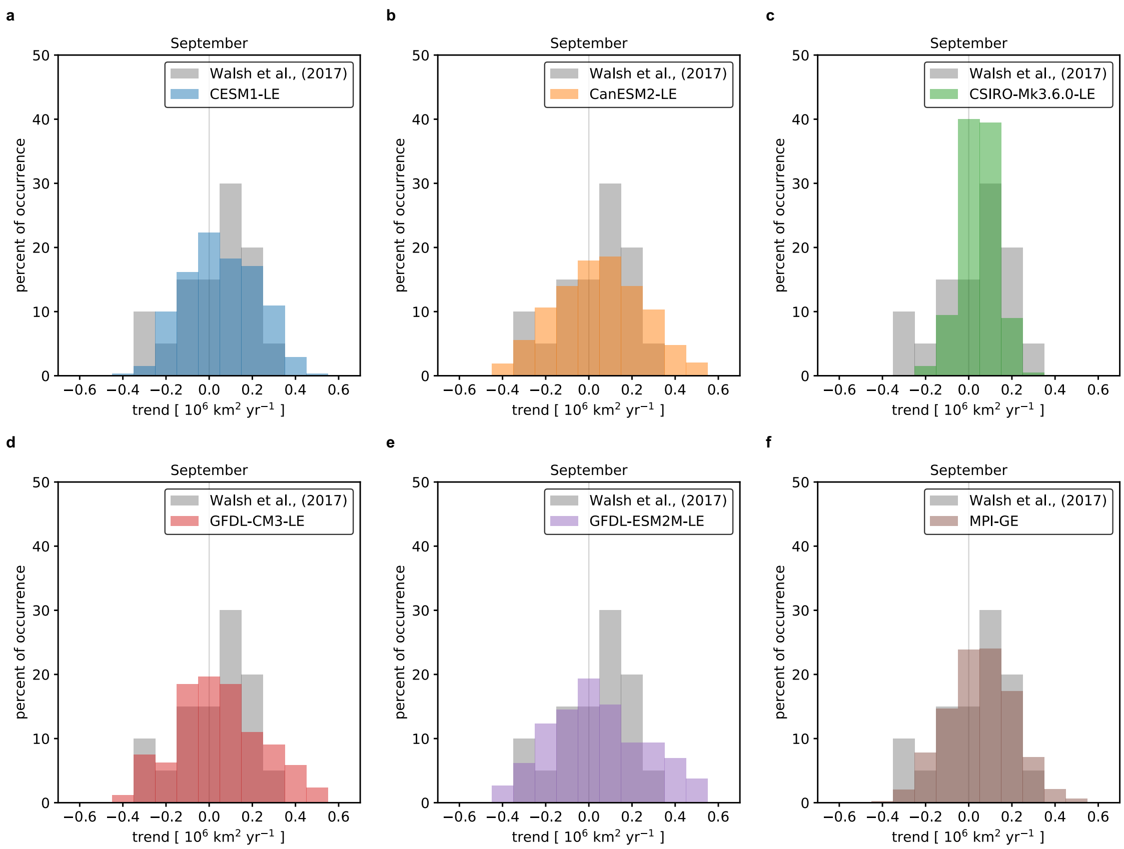

Figure 5. Percent of occurrence of separate 5-year trends in September Arctic sea-ice extent (SIE) from 1950-2019 for the (a) CESM1-LE, (b) CanESM2-LE, (c) CSIRO-Mk3.6.0-LE, (d) GFDL-CM3-LE, (e) GFDL-ESM2M-LE, and (f) MPI-GE. A 4th order polynomial was removed from each member of each SMILE prior to trend calculations to estimate the forced response. The bars show the distribution of trends for all members. The grey bars show percent of occurrence of separate 5-year trends in September Arctic SIE from 1930-2019 as estimated from Walsh et al. (2017). A 4th order polynomial was removed from the dataset prior to trend calculations to estimate the forced response.

4. Concluding remarks

The impacts of Arctic sea ice loss will be predominately felt by coastal communities, making it crucial to quantify and reduce projection uncertainty at regional scales. Here, we used a suite of SMILEs to investigate the sources of uncertainty in projections of Arctic SIE. For September SIE, model structure contributes between 40-80% of the total uncertainty over the next century, while for March SIE, model structure contributes approximately 40-90% of the total uncertainty over the next century and accounts for more uncertainty at the end of the 21st century. We find a clear timescale dependence for internal variability. For September SIE, internal variability contributes approximately 20-60% of total uncertainty in the next few decades, while for March SIE — and winter SIE more generally — internal variability contributes between 50-60% of total uncertainty and influences projections at longer lead times. Scenario uncertainty contributes mainly to uncertainty in summertime projections, accounting for approximately 70% of total uncertainty by the end of the century. We also find that the role for internal variability is mean-state dependent with thinner summer sea ice regimes more heavily influenced by internal variability, accounting for approximately 80% of total uncertainty for SIV $< 1{,}000$ km^3. At regional scales, the contribution of internal variability to total uncertainty increases, but has a large range and strongly depends on the month and region. In the GIN and Barents Seas, for instance, internal variability contributes approximately 50-70% of the total uncertainty over the next 30 years, while for the Central Arctic, internal variability accounts for approximately 20-30% of the total uncertainty. This is likely related to the influence of Atlantic heat transport on sea ice in the North Atlantic during the wintertime.

A unique result of our analysis is the partitioning of uncertainty due to different estimates of internal variability, which varies considerably across GCMs. This suggests that at least some GCMs are biased in their magnitude of variability. Due to the short observational record, it is difficult to precisely estimate the real-world magnitude of SIE internal variability (e.g., Brennan et al., 2020). However, using a reconstruction of September Arctic SIE to 1850 (Walsh et al., 2017) we try and estimate historical Arctic SIE variability. We limit our analysis to 1930 due to sparse data coverage in the Arctic prior to the 1930s. Figure 5 shows histograms of separate 5-year trends in September Arctic SIE from 1950-2019 using all members of each SMILE. A 4th order polynomial was used to approximate and remove the forced response to be consistent with comparison to observations. The grey bars indicate the range from Walsh et al. (2017) using separate 5-year trends from 1930 to 2019 after approximating the forced response as a 4th order polynomial fit and removing it. While most models appear to span the range of internal variability in the historical record, CSIRO-Mk3.6.0 does not simulate a large enough range of 5-year trends, most likely reflecting the fact that sea ice is biased high throughout the summer. This suggests the lowest contribution of internal variability to total uncertainty in projections September Arctic SIE is likely not realistic. Understanding and resolving these biases in internal variability across fully-coupled GCMs should remain a focus of the sea ice community as it is important for attribution of observed sea ice loss to anthropogenic climate change as well as for efforts of decadal prediction.

Recent work has highlighted the role of remote internal processes in determining sea ice trends across these same SMILEs (Topál et al., 2020), but a more process-oriented analysis of the spatial and temporal timescales of this variability may better reveal the sources of inter-model spread. For instance, it has been shown that these remote processes are not stable on longer time scales (Bonan and Blanchard-Wrigglesworth, 2020), suggesting that associated variability in September SIE during the satellite era does not paint a complete picture of the future SIE variability. The outsized role for internal variability in projections of Arctic sea ice changes in the coming decades further motivates the use of SMILEs to investigate a wide range of possible sequences of sea ice internal variability and its drivers. However, such work is beyond the scope of this paper, whose primary goal is to highlight the relative contribution of different sources of uncertainty to Arctic sea ice projections at different spatial and temporal scales.

While internal variability poses a great challenge for predicting Arctic SIE in the coming decades, the contribution of model structure to total uncertainty should not be ignored. So-called "emergent constraints", which link the inter-model spread in climate projections to observable predictors, should be used when characterizing projection uncertainty. Previous work has related the amount of future ice loss to the magnitude of historical SIE trends (Boé et al., 2009; Hall et al., 2019) and to the initial state of the sea ice (Bitz, 2008; Massonnet et al., 2012; Hall et al., 2019) and the Arctic climate (Senftleben et al., 2020), but open questions remain as to why these relationships exist and persist throughout the next century. Understanding biases in these trends (e.g., Rosenblum and Eisenman, 2016, 2017) and the physical mechanisms behind these constraints will improve the reliability of sea ice projections and increase confidence in our understanding of what controls Arctic sea ice loss.

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