

## Supplementary Material

J. J. Buffo<sup>1</sup>, B. E. Schmidt<sup>2</sup>, C. Huber<sup>3</sup>, C. R. Meyer<sup>1</sup>

<sup>1</sup>Dartmouth College, <sup>2</sup>Georgia Institute of Technology, <sup>3</sup>Brown University

Contained below are three supplementary material sections which support and/or provide additional detail to the work and conclusions presented in the main manuscript: *Characterizing the Ice-Ocean Interface of Icy Worlds: A Theoretical Approach*. Section S1 provides the complete solution to the molecular diffusion equation in the mushy layer – as outlined in Section 2.1.1 of the main text. Section S2 provides a table of all variables used throughout the text as well as their associated symbols and units. Finally, Section S3 provides a table of the ionic species present and their relative abundances in the terrestrial ocean composition we assume throughout the investigation.

## S1. Solving the Molecular Diffusion Equation

The evolution of salinity at the ice-mush interface is governed by:

$$\frac{\partial S}{\partial t} = D \frac{\partial^2 S}{\partial z^2} \quad (S1)$$

Utilizing the similarity variable  $\eta = z/2\sqrt{Dt}$  and boundary conditions (Equations 9-10) it can be shown that:

$$\theta = \frac{1}{\text{erfc}(\lambda_S)} - \frac{\text{erf}(\eta)}{\text{erfc}(\lambda_S)} \quad (S2)$$

Writing in terms of the original variables  $S$ ,  $z$  and  $t$ :

$$S = S_{oc} + (S_{int} - S_{oc}) \left( \frac{1}{\text{erfc}(\lambda_S)} - \frac{\text{erf}(z/2\sqrt{Dt})}{\text{erfc}(\lambda_S)} \right) \quad (S3)$$

The Stefan condition for this problem can be garnered from the equation for conservation of salt (no salt in the ice phase):

$$\int_{z_m^*(t)}^{\infty} S(z, t) dz = \text{cnst.} \quad (S4)$$

Taking the temporal derivative of this equation and applying the Leibniz integral rule:

$$-S(z_m^*, t) \frac{dz_m^*}{dt} + \int_{z_m^*(t)}^{\infty} \frac{\partial S}{\partial t} dz = 0 \quad (S5)$$

Substituting  $\frac{\partial S}{\partial t} = D \frac{\partial^2 S}{\partial z^2}$ , noting  $S(z_m^*, t) = S_{int}$  and carrying out the integral gives:

$$S_{int} \frac{dz_m^*}{dt} = -D \frac{\partial S(z = z_m^*)}{\partial z} \quad (S6)$$

The derivatives are:

$$\frac{dz_m^*}{dt} = \frac{\lambda_S \sqrt{D}}{\sqrt{t}} \quad (S7)$$

And

$$\frac{\partial S(z = z_m^*)}{\partial z} = \frac{-(S_{int} - S_{oc}) \exp(-\lambda_S^2)}{\text{erfc}(\lambda_S) \sqrt{\pi Dt}} \quad (S8)$$

Substituting back into Equation 17,

$$S_{int}\sqrt{\pi} = \frac{(S_{int} - S_{oc})}{\lambda_S \operatorname{erfc}(\lambda_S)} \exp(-\lambda_S^2) \quad (S9)$$

Or, rearranging

$$\lambda_S \operatorname{erfc}(\lambda_S) \exp(\lambda_S^2) = \frac{(S_{int} - S_{oc})}{S_{int}\sqrt{\pi}} \quad (S10)$$

We assume that the interface is at its melting temperature,  $T_f$ , which we take as a linear function of salinity  $T_f = T_{mp} - 0.066178S$ , where  $T_{mp}$  is the melting temperature of pure ice and the freezing point depression coefficient has units of K kg g<sup>-1</sup>. Solving for  $S$  and letting  $T_f$  lie on a conductive (linear) thermal profile  $T(z) = T_S + z(T_{oc} - T_S)/H$  at a depth  $H-h$ . We have:

$$S_{int} = 15.1106(T_{mp} - T_f) = 15.1106 \left( T_{mp} - \left( T_S + (H-h) \frac{(T_{oc} - T_S)}{H} \right) \right) \quad (S11)$$

Substituting the value of  $S_{int}$  into Equation 21 and setting the interface velocities equal ( $\dot{z}_m^* = 2\lambda_S\sqrt{D}/\sqrt{t} = 2\lambda_T\sqrt{\kappa_i}/\sqrt{t} = \dot{z}_m$ ) produces two equations:

$$\lambda_S = \frac{\lambda_T\sqrt{\kappa_i}}{\sqrt{D}} \quad (S12)$$

and

$$15.1106 \left( T_{mp} - \left( T_S + (H-h) \frac{(T_{oc} - T_S)}{H} \right) \right) \sqrt{\pi} = \frac{15.1106 \left( T_{mp} - \left( T_S + (H-h) \frac{(T_{oc} - T_S)}{H} \right) \right) - S_{oc}}{\lambda_S \operatorname{erfc}(\lambda_S)} \exp(-\lambda_S^2) \quad (S13)$$

which can be solved for  $h$ . The equation is linear in  $h$ , so has one solution.

Archie's law is employed to estimate the molecular diffusion in a porous medium and it is assumed that transport processes (diffusion and advection) will be limited by the critical porosity of ocean/brine derived ices ( $\phi_c = 0.05$  [K M Golden *et al.*, 2007]):

$$D = k_S \phi_c^m \quad (S14)$$

Where  $k_S$  is the molecular diffusivity of salt in water and  $m$  is a cementation exponent, here=2, which describes how ion transport is limited by porosity.

## S2. Variables Used in the Text

Symbol	Definition	Units
$\alpha$	1D Advection Coefficient	$\text{kg m}^{-3} \text{s}^{-1}$
$\beta$	Density (Salinity) Coefficient	$\text{ppt}^{-1}$
$br^{\uparrow, \downarrow}$	Vertical Brine Velocity	$\text{m s}^{-1}$
$c_i$	Ice Heat Capacity	$\text{J kg}^{-1} \text{K}^{-1}$
$D$	Salt Diffusivity	$\text{m}^2 \text{s}^{-1}$
$\eta$	Similarity Variable	-
$\eta_m$	Similarity Variable at $z_m^*$	-
$\Gamma$	Freezing Point Depression Coefficient	$\text{K kg g}^{-1}$
$g$	Acceleration Due to Gravity	$\text{m s}^{-2}$
$G$	Thermal Gradient	$\text{K m}^{-1}$
$G'$	Thermal Gradient Perturbation	$\text{K m}^{-1}$
$h$	Mushy Layer Equilibrium Thickness	m
$h'$	Thickness Perturbation	m
$H$	Ice Shell Thickness	m
$k_s$	Salt Diffusivity in Pure Water	$\text{m}^2 \text{s}^{-1}$
$\kappa_i$	Thermal Diffusivity of Ice	$\text{m}^2 \text{s}^{-1}$
$\kappa_{br}$	Thermal Diffusivity of Brine	$\text{m}^2 \text{s}^{-1}$
$L$	Latent Heat of Fusion	$\text{J kg}^{-1}$
$Le$	Lewis Number	-
$\lambda_T, \lambda_S$	Stefan Problem Variables	-
$m$	Cementation Exponent	-
$\mu$	Kinematic Viscosity	$\text{m}^2 \text{s}^{-1}$
$\phi$	Liquid Fraction	-
$\phi_c$	Critical Porosity	-
$\Pi$	Permeability	$\text{m}^2$
$Ra_c$	Critical Rayleigh Number	-
$\rho_{sw}$	Ocean Density	$\text{kg m}^{-3}$
$S$	Salinity	ppt
$S_{int}$	Interface Salinity	ppt
$S_{oc}$	Ocean Salinity	ppt
$\Delta S_j$	$S_{int} - S_{oc}$	ppt
$t$	Time	s
$T$	Temperature	K
$T_m$	Melting/Freezing Temperature	K
$T_{oc}$	Ocean Temperature	K
$T_s$	Surface Temperature	K
$\theta$	Dimensionless Salinity	-
$v_m$	Mush-Ocean Freezing Front Velocity	$\text{m s}^{-1}$
$v_m^*$	Ice-Mush Freezing Front Velocity	$\text{m s}^{-1}$
$\Delta V$	$v_m - v_m^*$	$\text{m s}^{-1}$
$z$	Vertical Coordinate	m
$z_m$	Mush-Ocean Interface	m
$z_m^*$	Ice-Mush Interface	m

### S3. Terrestrial Ocean Composition

Species	Terrestrial Seawater (mol/kg)
Na <sup>+</sup>	4.69 x 10 <sup>-1</sup>
K <sup>+</sup>	1.02 x 10 <sup>-2</sup>
Ca <sup>2+</sup>	1.03 x 10 <sup>-2</sup>
Mg <sup>2+</sup>	5.28 x 10 <sup>-2</sup>
Cl <sup>-</sup>	5.46 x 10 <sup>-1</sup>
SO <sub>4</sub> <sup>2-</sup>	2.82 x 10 <sup>-2</sup>
<b>Total Salt (ppt)</b>	34

**Table S1 – Ocean composition.** List of ion species and relative abundances for terrestrial seawater [*Dickson and Goyet, 1994*].